

NZ Centre for Earthquake Resilience

# Insights from ground-motion simulation validation of New Zealand small magnitude subduction earthquakes

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# **1.** Unique characteristics warrant special treatment

Physics-based ground-motion simulation can provide predictions of crustal earthquake ground-motion intensity metrics in New Zealand which are better than empirical ground-motion models. This study aims to extend this level of performance to subduction interface and slab earthquakes which have unique source rupture and travel path characteristics:

## Interface earthquakes:

## **3.** Models for rupture characteristics

To gain insights on systematic differences between crustal, interface, and slab earthquake source ruptures, the models contained in SRCMOD (Mai et al. 2014), an online catalogue of finite-fault rupture models, were analysed. The variations of the betweenevent residuals with magnitude and hypocentre depth for these three tectonic classifications are shown in Figure 4–5. By comparing between-event residuals of risetime and relative rupture velocity and the dependence on hypocentre depth and magnitude, it was possible to infer systematic differences between crustal, interface, and slab ruptures.

# 5. Simulation validation

The new subduction models were implemented within the hybrid broadband ground-motion simulation approach developed by Graves and Pitarka 2010, 2015, 2016. Ground-motion predictions were validated using high-quality small magnitude subduction earthquake ground-motions in New Zealand. The geospatial distribution of earthquakes and stations used for validation are shown in Figure 6—only earthquakes and stations with atleast three associated high-quality observed ground-motion records were considered.

- Rupture interface at interplate boundary • Effect of subducted sediment and seamounts • Potential for very large ruptures (Figure 1–3) Slab earthquakes:
- Deep ruptures within descending subducted slab • Occur in high-stress/temperature environments
- Effect of volcanic arc on energy attenuation

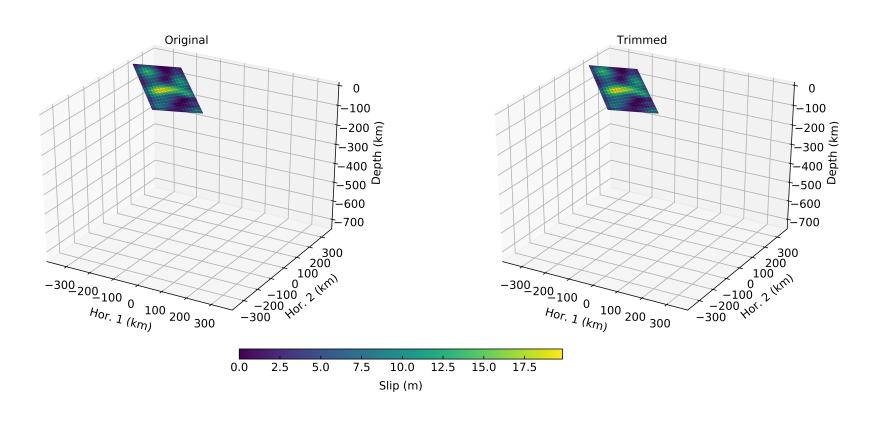
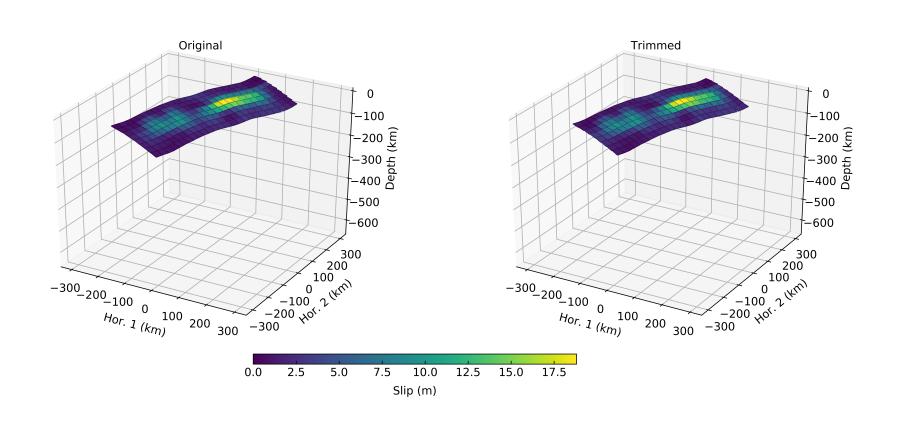


Figure 1: Rupture model for 2004  $M_w$ 9.2 Sumatra earthquake (Ji 2005)



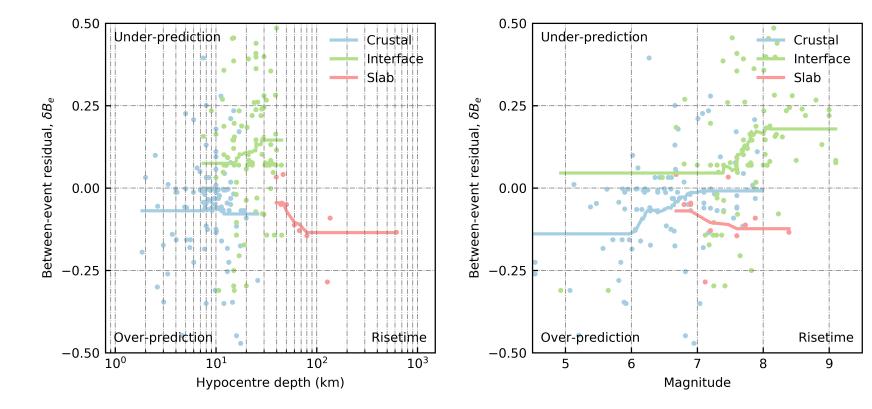


Figure 4: Between-event residuals,  $\delta B_e$ , for risetime

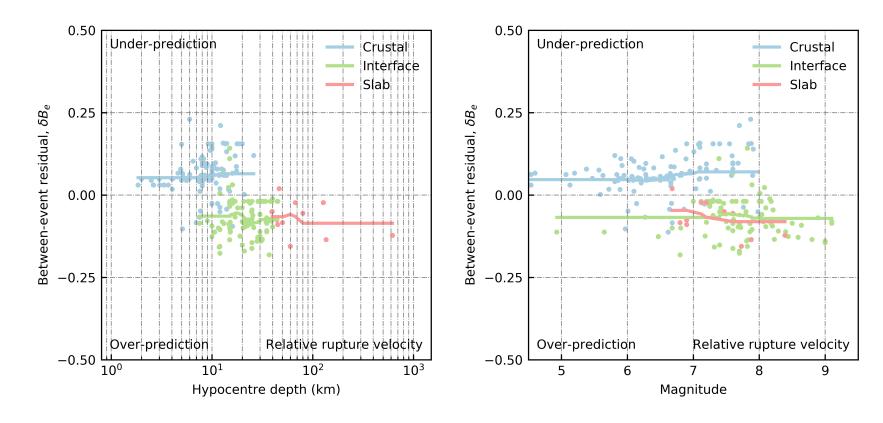


Figure 5: Between-event residuals,  $\delta B_e$ , for rupture velocity

New models for subduction earthquake ruptures were developed based on the observed trend for residuals of risetime and relative rupture velocity:

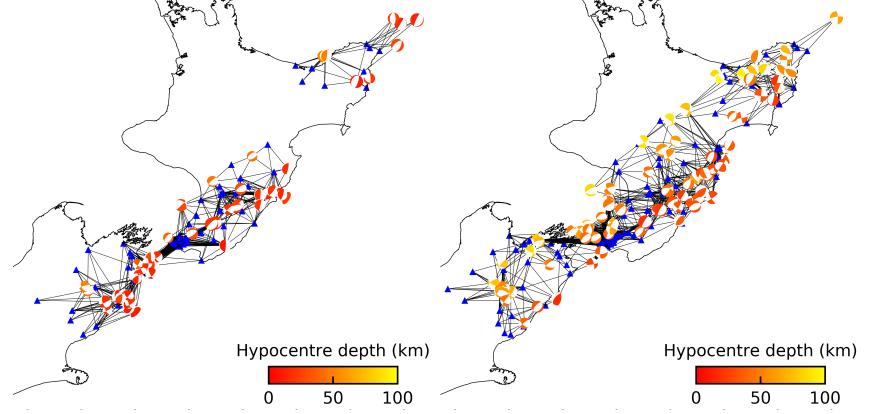


Figure 6: Validation data set; left: interface earthquakes; right: slab earthquakes

## 6. Results and next steps

The predictions for subduction earthquakes done with the new subduction simulation models show that the simulations are performing well for subduction earthquakes. The simulations are able to provide predictions for small magnitude subduction earthquakes in New Zealand which are more accurate than global empirical ground-motion models for subduction earthquakes. The simulations are now able to provide comparable levels of predictive accuracy for small magnitude subduction and crustal earthquakes in New

#### Figure 2: Rupture model for 2010 $M_w 8.8$ Maule earthquake (Lorito et al. 2011)

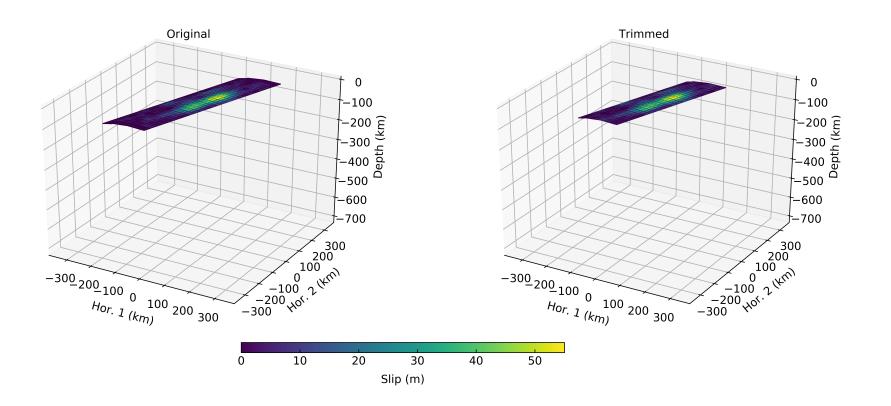


Figure 3: Rupture model for 2011  $M_w$ 9.0 Tohoku earthquake (Hayes 2013)

## 2. Methodology

A mixed-effects regression approach, in which residuals are partitioned based on causative effects was used in several portions of this study:

• Analysis of a catalogue of finite-fault models to inform models for subduction ruptures

$$\Delta \sigma_{if} = 25 + 0.5 \times depth \tag{2}$$

$$\Delta \sigma_{slab} = 50 + 1.5 \times depth \tag{3}$$

$$Rup.velocity_{if} = 70\% \times \beta \tag{4}$$

$$Rup.velocity_{slab} = 90\% \times \beta \tag{5}$$

where  $\beta$  is the shear wave velocity at the source and  $\Delta \sigma$  is stress parameter. (*Rup.velocity*<sub>crstl</sub> = 80\% ×  $\beta$ ,  $\Delta \sigma_{crstl} = 50 \text{ bar}$ 

## 4. Models for path effects

To account for greater backarc anelastic attenuation determined from analysis of empirical ground-motion models, a record-specific adjustment factor,  $\phi$  was computed based on the absolute value of the source-tosite azimuth from the backarc,  $\theta$ , and a source-specific adjustment factor  $\eta$  was computed based on the source hypocentre depth, where:

#### Zealand.

Investigation of the spatial distribution of betweenevent and systematic site-to-site residuals indicates that some regional trends in prediction misfit persist (Figure 7 and 8). In particular, there appears to be over-prediction for sites located along the forearc and for deep slab events in the Taupō volcanic zone.

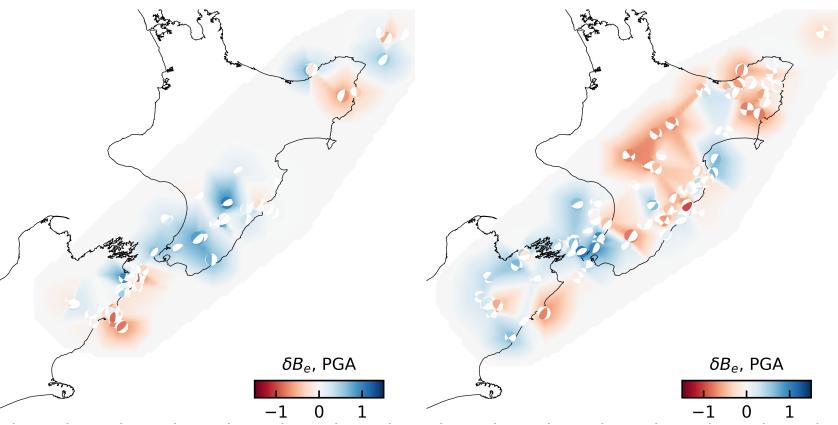
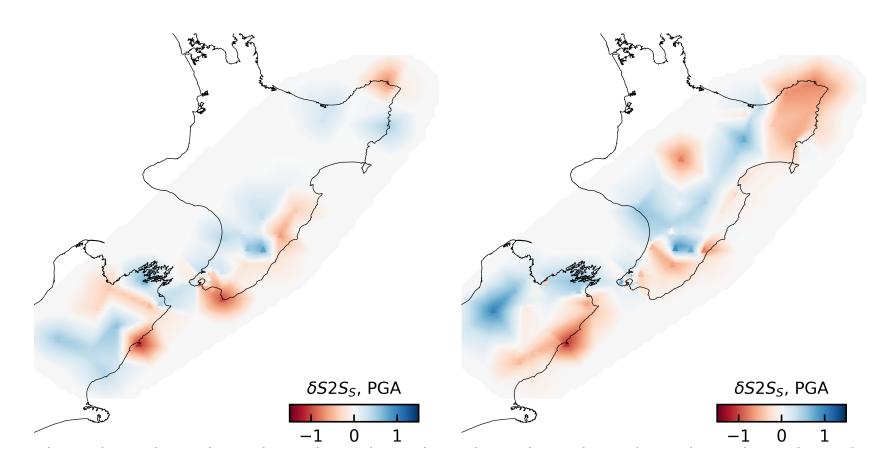


Figure 7: Between-event residuals,  $\delta B_e$ , for peak ground acceleration (PGA); left: interface ground-motions; right: slab groundmotions



- Analysis of empirical ground-motion model predictions to inform models for volcanic arc path effects • Validation of simulation predictions for subduction
- earthquake ground-motions using the new models

The general form of the equation is:

 $\ln IM_{es} = f_{es} + a + \delta B_e + \delta W_{es}$ (1)

where  $\ln IM_{es}$  is the natural logarithm of the reference intensity metric (IM);  $f_{es}$  is the median of the predicted logarithmic IM for event, e, and site, s, either from a simulation or empirical GMM; a is the predictive model bias;  $\delta B_e$  is the between-event residual with zero mean and variance  $\tau^2$ ; and  $\delta W_{es}$  is the within-event residual with zero mean and variance  $\phi^2$ .

$$\phi_{if,slab} = \begin{cases} -0.3, & \text{for } 0 \leq \theta \leq 80 \\ -0.3 + 0.03(\theta - 80), & \text{for } 80 \leq \theta \leq 100 \\ 0.3, & \text{for } 100 \leq \theta \leq 180 \\ (6) \end{cases}$$

$$\eta_{if,slab} = \begin{cases} 0, & \text{for } 0 \leq depth \leq 40 \\ (depth - 40)/60, & \text{for } 40 \leq depth \leq 100 \\ 1, & \text{for } 100 \leq depth \end{cases}$$

$$(7)$$
These adjustment factors were used to modify the rock-quality factors  $Q_P$  and  $Q_S$ , which control anelastic attenuation in the HF simulation component:

 $Q_{if.slab} = Q \times (1 + \phi_{if.slab}\eta_{if,slab})$ (8)

Figure 8: Systematic site-to-site residuals (the portion of  $\delta W_{es}$ ) ascribed to systematic site effects),  $S2S_s$ , for peak ground acceleration (PGA); left: interface ground-motions; right: slab ground-motions

## Future work will focus on:

- Reduction of the observed regional residual trends with tuning of the volcanic arc-based modifications to the rock-quality factors  $Q_P$  and  $Q_S$
- Extend the models validated on small magnitudes to large magnitude and megathrust scenario ruptures of subduction earthquakes